

## 12.843 - project 2

### Derivations

#### *Shear layer*

1a) Shear layer instability: consider

$$q = \zeta_0 + \Delta_1 \mathcal{H}(y - y_1 - \eta_1(x, t)) + \Delta_2 \mathcal{H}(y - y_2 - \eta_2(x, t))$$

Let  $y_1 = -W/2$  and  $y_2 = W/2$ . Linearize, assuming  $\eta_j = n_j \exp(i k [x - ct])$  and  $n_j$  is suitably small. Write the matrix eigenvalue equation for the  $n$ 's.

1b) Rayleigh theorem: Show that instability requires  $\Delta_1 \Delta_2 < 0$ . Hint: consider

$$\begin{pmatrix} m_{11} & m_{12} \\ m_{21} & m_{22} \end{pmatrix} \begin{pmatrix} n_1 \\ n_2 \end{pmatrix} = c \begin{pmatrix} n_1 \\ n_2 \end{pmatrix}$$

and figure out the constraints on  $m_{21}m_{12}$  such that the imaginary part of  $c$  could be nonzero.

1c) Fjortoft theorem: Show that you stabilize the flow when  $U_i \Delta_i$  becomes negative.

1d) Plot  $c$  for  $kW$  from near zero to 3. Plot the growth rate  $k \operatorname{Im}(c)$ .

#### *Adjustment*

2a) Consider an initial state

$$v = A \exp(-k|x|) \quad , \quad h = H \quad , \quad u = 0$$

From the linearized SW dynamics, find the final state assuming all the gravity waves radiate away. Evaluate the ratio of the geostrophic final state energy to the initial energy as a function of  $kR_d$ .

2b) Contrast this with the case

$$v = 0 \quad , \quad h = H + A \exp(-k|x|) \quad , \quad u = 0$$

2c) consider what might be a simple, qualitative way to understand the differences: how large a region is out of balance initially for the two cases and small and large  $k$  values? This indicates how much mass needs to be moved and how much energy might be left for the steady state.

## Numerical labs

### *Frontal waves and shear instability*

From the derivation with  $\zeta_0 = 0$ , the flow is unstable for  $kW$  smaller than 1.2785. Verify this is a good estimate from your analysis. Show that the numerical model (“shear layer”) with  $k = 1$ , gives reasonable agreement with this (e.g.,  $W = 1.25$  is unstable, 1.3 is not). But we now want to look at the question of nonlinear stability: is, in fact,  $W = 1.3$  stable as we increase the perturbation amplitude? A related question is whether the unstable mode just below the critical condition equilibrates at some amplitude which gets smaller as  $kW$  approaches the value above?

I’ve also put up a code which computes the stability for a tanh shear layer on a  $\beta$ -plane. For  $U = U_0 \tanh(y/W)$ , what value of  $\beta$  will stabilize the flow?

You can use the code (with  $\beta = 0$ ) to see what happens as  $W$  becomes less than the critical width, which, in this case is 1. Is the nonlinear evolution similar or different from the broken line profile? What happens for  $W = 0.5$  as you increase  $\beta$ ? Use the stability code to estimate the value for which the flow becomes stable. It should be

$$\frac{\beta W^2}{U_0} = \frac{4}{3\sqrt{3}}$$

with  $U_0$  the maximum basic state velocity (1 in the code). How is it different with discrete wavenumbers?

### *Adjustment*

Using the 2D SW code and inversion codes, examine the behavior of a dipolar initial condition

$$Q = \epsilon x \exp(-[x^2 + y^2]/2)$$

Try and monitor gravity wave energy by looking at a few points away from the vortex. Consider

2a) unbalanced initial conditions [ $h' = -Q/\gamma^2(1 + Q)$ ,  $\psi = 0$ ]

2b) geostrophically balanced [ $h' = \psi$ ]

2c) nonlinear balance

What happens for small and medium Rossby numbers (less than one, though)?